Application of Conservation Laws with Source Terms to the Shallow Water Equations and Crowd Dynamics

Polly Smith

September 2004

Abstract

In this dissertation we look at the application of two dimensional conservation laws with source terms to two specific physical examples, the shallow water equations and crowd flow. There are a number of techniques that can be used to numerically approximate such systems. We discuss the Q scheme method of Bermúdez & Vázquez, focusing on an extension to Roe's Q scheme. The technique of dimensional splitting is used to decompose the systems into two one-dimensional problems that are solved alternately in the x and y directions. For the shallow water equations we examine two di erent test problems. We begin by considering a non-rotating frame of reference, before extending the equations to include the e ects of the Earth's rotation. The ¯-plane approximation is used to avoid the complications of spherical geometry whilst retaining the leading order e ect of the Earth's curvature. We then move on to develop a two dimensional macroscopic model of crowd flow based on a one-dimensional continuum model proposed by Payne and Whitham. The equations are transformed into a system of conservation laws, with source term in the form of a relaxation term. A series of numerical experiments are used to evaluate the accuracy of the model.

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Declaration

I confirm that this is my own work and the use of all materials from other sources has been properly and fully acknowledged.

Signed

Polly Smith September 2004

Acknowledgements

Firstly, I would like to thank my supervisor Professor Mike Baines for his continual help and support throughout this project. His patience and humour is admirable. I am also very grateful to the postgraduate secretary Sue Davis and my fellow MSc students for their support and encouragement over the past year.

Apologies go to my family for all the angst - thank you for your tolerance. A big thank you to all my friends for always being there despite my neglect, and to Tim for listening to my worries.

Many thanks to the NERC for their generous financial support.

Contents

Chapter 1

Introduction

Hyperbolic systems of conservation laws with source term arise naturally in a wide variety of applications. Many practical problems in science and engineering involve mathematical models based on this type of system. In two space dimensions the equations take the form

$$
\frac{\mathcal{Q}_W}{\mathcal{Q}_t} + \frac{\mathcal{Q}F(w)}{\mathcal{Q}_X} + \frac{\mathcal{Q}G(w)}{\mathcal{Q}_Y} = R
$$

where w is a vector of conserved quantities, $F(w)$ and $G(w)$ are functions representing the physical flux, and \bf{R} is a source term. Often, the flux functions are nonlinear functions of w, leading to nonlinear systems of partial di erential equations (PDEs). Some of these systems have exact solutions, but generally analytical solutions to these problems are not available since they are extremely dicult, if not impossible, to find. Thus we seek to devise numerical methods for their approximate solution. Indeed, the significance of these equations to a number of important fields has provoked increasing research in recent years.

In this project we look at the application of conservation laws with source

term in two dimensions by considering two physical examples - The shallow water equations and crowd dynamics. Many similarities exist between the flow of crowds and that of a shallow fluid. The movement of a large crowd exhibits many of the attributes of a fluid and macroscopic pedestrian models utilise this analogy by treating the crowd as a continuum. There are, however, drawbacks to this assumption as unlike a fluid, a crowd has the ability to think and will react to it surroundings.

Chapter 2 introduces the numerical method that will be used to solve the two dimensional hyperbolic systems of equations arising from the shallow water and crowd models. In chapter 3 we derive the shallow water equations for two dimensional, incompressible, irrotational flow in a rectangular channel and apply the approach discussed in Chapter 2 to the model. We then extend the model to include the e ects of the Earth's rotation by introducing

Chapter 2

Numerical Methods

There are a number of techniques that can be used to numerically approximate the two dimensional system of conservation laws with source term

$$
\frac{\mathscr{E}w}{\mathscr{E}t} + \frac{\mathscr{E}F(w)}{\mathscr{E}x} + \frac{\mathscr{E}G(w)}{\mathscr{E}y} = R \tag{2.1}
$$

such as finite element, finite volume and finite di erence methods. In this chapter we use the approach of dimensional splitting to derive a second order accurate scheme that can be used to solve systems of the type (2.1). We then introduce the Q Schemes of Bermúdez & Vázquez [1] and explain how they can be used to give expressions for the numerical flux and numerical source terms. Finally, we consider the specific example of Roe's Q Scheme.

2.1 Dimensional Splitting

The method of dimensional splitting was first proposed by Yanenko [20].

By rewriting the source term as

$$
R=f+g\H;
$$

directions respectively and $\mathsf{F}^*_{i+\frac{1}{2}}$ and $\mathsf{G}^*_{i,j+\frac{1}{2}}$ are the corresponding numerical flux functions. $\mathbf{R}_{i:j}^*$ is a numerical source function taken to be of the form

$$
\mathsf{R}_{i;j}^* = \mathsf{f}_{i;j}^* + \mathsf{g}_{i;j}^*
$$

where ${\sf f}_{i;j}^*$ and ${\sf g}_{i;j}^*$ are the source term approximations containing the x and y direction components only.

The boundaries are given by the lines $(x_{-\frac{1}{2}}; y_j)$, $(x_{M+\frac{1}{2}}; y_j)$, $(x_i; y_{-\frac{1}{2}})$ and $(x_i, y_{N+\frac{1}{2}})$ where M and N denote the number of intervals in the x and y directions respectively. Numerical boundary conditions are required at each of these points and will be discussed in more detail later both for the shallow water equations and crowds.

The simplest form of splitting then discretises in two steps. First, we solve in the x coordinate direction using a half time step $\frac{\Delta t}{2}$ to obtain an approximation of (2.2), $w^{n+\frac{1}{2}}$,

$$
w_{ij}^{n+\frac{1}{2}} = w_{ij}^{n} i \frac{t}{2-x}^{3} F_{i+\frac{1}{2};j}^{*} i F_{i-\frac{1}{2};j}^{*}
$$

initial conditions to obtain w^{n+1} .

$$
w_{ij}^{n+\frac{1}{4}} = w_{ij}^{n} i \frac{t}{4} \overline{x}^{3} F_{i+\frac{1}{2}j}^{*} i F_{i-\frac{1}{2}j}^{*} + \frac{t}{4} \overline{x}^{f}_{ij}^{*}
$$

\n
$$
w_{ij}^{n+\frac{3}{4}} = w_{ij}^{(n+\frac{1}{4})} i \frac{t}{2} \overline{y}^{3} G_{ij+\frac{1}{4}}^{*(n+\frac{1}{4})} i G_{ij-\frac{1}{2}}^{*(n+\frac{1}{4})} + \frac{t}{2} \overline{y}^{*(n+\frac{1}{4})}
$$

\n
$$
w_{ij}^{n+1} = w_{ij}^{(n+\frac{3}{4})} i \frac{t}{4} \overline{x}^{*(n+\frac{3}{4})} i F_{i+\frac{1}{2}j}^{*(n+\frac{3}{4})} i F_{i-\frac{1}{2}j}^{*(n+\frac{3}{4})} + \frac{t}{4} f_{ij}^{*(n+\frac{3}{4})}.
$$

2.1.1 Stability

A necessary condition for the stability of a homogeneous system of equations

$$
\frac{\mathscr{E} \mathbf{w}}{\mathscr{E} t} + \frac{\mathscr{E} \mathsf{F}(\mathbf{w})}{\mathscr{E} \mathsf{x}} = 0;
$$

is

$$
max_{i,j,k} F \frac{t}{x} j \qquad \circ \qquad (2.4)
$$

where \Box^F_k are the eigenvalues of the Jacobian matrix ${\mathbf A}$ and \circ is a bound constant known as the Courant or CFL number (after Courant, Friedrichs and Lewy).

In this instance we have two stability conditions, (2.4) for the x direction and

$$
max_j j_s g \frac{t}{y} j \qquad c.
$$
 (2.5)

for the y direction. To ensure that both (2.4) and (2.5) are satisfied and that the dimensional splitting scheme remains stable we choose our time step t such that

$$
t \cdot \frac{min(\mathbf{x}; \mathbf{y})}{max_{i,j}(\mathbf{y}, \mathbf{f}; \mathbf{y}, \mathbf{G})} \quad :
$$

2.2 Q Schemes

In this section we discuss a numerical technique proposed by Bermúdez & Vázquez [1] for approximating two dimensional hyperbolic systems of conservation laws with source term.

In [1] Bermúdez & Vázquez introduce a general class of upwind methods for solving nonhomogeneous problems. They propose extensions to a number of well known and well studied flux di erence and flux splitting techniques using an upwinded discretisation of the source term. In particular, a class of upwind schemes known as Q schemes are considered. The schemes were

The matrix Q depends continuously on w_L and w_R and is characteristic of each Q scheme.

Bermúdez & Vázquez [1] propose upwinding the source term in a similar manner to the flux. The numerical source functions f^{*} and g^{*} are taken to be of the form

$$
\mathbf{f}_{i,j}^* = \mathbf{f}_{i+\frac{1}{2},j}^* + \mathbf{f}_{i-\frac{1}{2},j}^* \text{ and } \mathbf{g}_{i,j}^* = \mathbf{g}_{i,j+\frac{1}{2}}^* + \mathbf{g}_{i,j-\frac{1}{2}}^* \tag{2.8}
$$

The left and right numerical sources ${\sf f}^*_{i-\frac{1}{2},j'}$, ${\sf f}^*_{i+\frac{1}{2},j}$ are obtained by projecting the source term onto the eigenvectors of Q to give

$$
\mathbf{f}_{i-\frac{1}{2};j}^{*} = \frac{1}{2} \mathbf{f}_{i} + jQjQ^{-1} \mathbf{G}_{i-\frac{1}{2};j} \hat{\mathbf{f}}_{i-\frac{1}{2};j}
$$
(2.9)

and

$$
\mathbf{f}_{i+\frac{1}{2};j}^{*} = \frac{1}{2} \mathbf{\hat{L}} + j \mathbf{Q} j \mathbf{Q}^{-1} \mathbf{I}_{i+\frac{1}{2};j} \hat{\mathbf{f}}_{i+\frac{1}{2};j}
$$
(2.10)

where \hat{f} represents an approximation of f . Expressions for g^* are obtained by symmetry.

 jQj can be obtained by using

$$
jQj = X^{-1}j \, jX
$$

where represents a matrix whose diagonal elements are the eigenvalues of

Q, and X is a matrix containing the right eigenvectors.

Note that since

$$
f^*(w_L; w_R)
$$
 j! $f(w)$ as $w_L; w_R$ j! w

we have consistency of the numerical source with the source term.

2.2.1 Roe's Q scheme

Roe's Q scheme [13] takes ${\sf Q}_{i+\frac{1}{2},j}$ and ${\sf Q}_{i:j+\frac{1}{2}}$ equal to the Roe averaged Jacobian matrices \tilde{A} and \tilde{B} , i.e.

$$
Q_{i+\frac{1}{2},j}^n = Q^{\dagger} w_{ij}^n / w_{i+1,j}^n{}^{\mathbb{C}} = \tilde{A} \text{ and } Q_{ij+\frac{1}{2}}^n = Q^{\dagger} w_{ij}^n / w_{ij+1}^n{}^{\mathbb{C}} = \tilde{B}.
$$

where

$$
\tilde{A} = A(\tilde{w}) \mathcal{U} \frac{\partial F}{\partial w} \quad \text{and} \quad \tilde{B} = B(\tilde{w}) \mathcal{U} \frac{\partial G}{\partial w}
$$

are locally linearised Jacobian matrices. The term \tilde{w} is known as the Roe average of w_L and w_R , where w_L , w_R represent adjacent piecewise constant states. \tilde{A} , \tilde{B} and therefore Q are chosen so that the following properties are satisfied

$$
^2
$$
 $Q(w_L; w_R)$ is diagonalisable with real eigenvalues (hyperbolicity);

$$
^2
$$
 $Q(w_L; w_R)$ *i* P $Q(w)$ as $w_L; w_R$ *i* P w (consistency):

$$
P = F(w_{L,j}) + F(w_{R,j}) = Q(w_{L,j}; w_{R,j}) (w_{L,j} + w_{R,j})
$$
 and:

$$
^2 \mathbf{G}(\mathbf{w}_{i;L}) \mathbf{j} \mathbf{G}(\mathbf{w}_{i;R}) = \mathbf{Q}(\mathbf{w}_{i;L}; \mathbf{w}_{i;R}) (\mathbf{w}_{i;L} \mathbf{j} \mathbf{w}_{i;R})
$$
 (conservation).

The final property ensures that we obtain a conservative algorithm and that any discontinuity moves at the correct speed.

By discretising the system (2.1) separately in the two coordinate directions we obtain expressions for \tilde{w} and Q along the x and y coordinate lines respectively. Equations (2.6) - (2.10) can then be used to give the two dimensional numerical flux and numerical source functions.

Chapter 3

The Shallow Water Equations

Meteorology and weather prediction are both fields in which conservation laws apply. The equations used to describe atmosphere and ocean processes are often too complex to solve analytically. In this case numerical methods are applied to obtain an approximate solution.

The shallow water equations can be used to simulate the flow of water in rivers, lakes and coastal regions. Thus they provide a valuable means for modelling a variety of problems related to the environment and coastal engineering etc. For instance, sediment transport in coastal regions has recently become a major topic of interest in the hydraulics industry [9].

3.1 The Equations

In this section we introduce the equations governing the flow of an incompressible fluid in a shallow domain, i.e. a domain whose depth is small relative to the dimensions of the problem. Initially, we will consider the shallow water approximation in the case of irrotational flow before introducing the e ect

of the fluid at the same point but from a fixed reference level, and g is the gravitational constant, see figure 3.1. It is often convenient to rewrite $H(x, y)$ in terms of the bed height $B(x; y)$

$$
H(x; y) = D_j B(x; y)
$$

3.2 Conservation form

To rewrite the system (3.1) - (3.3) in generalised conservation form we first multiply (3.1) by u_i

$$
uh_t + u(uh)_x + u(hv)_y = 0. \tag{3.4}
$$

Using the product rule

$$
(uh)_t = hu_t + uh_t;
$$

to substitute in for hu_t after multiplying (3.2) by h gives

$$
(uh)_{t i} uh_t + (hu)u_x + (hv)u_y + ghh_x = ghH_x.
$$
 (3.5)

Using (3.4) and the product rule on $(hu^2)_{x}$, $(huv)_y$ and $(gh^2)_{x}$, i.e.

$$
(huu)_x = (hu)_x u + (hu)u_x
$$

$$
(huv)_y = (hv)_y u + (hv)u_y
$$

$$
(gh^2)_x = 2ghh_x
$$

and substituting into (3.5) we obtain (3.2) with the left hand side in conservative form

$$
(uh)_t + (u^2h + \frac{1}{2}gh^2)_x + (huv)_y = ghH_x.
$$

Following a similar procedure for (3.3) yields

$$
(vh)_t + (huv)_x + (v^2h + \frac{1}{2}gh^2)_y = ghH_y
$$

We can now write the system in vector form as

$$
\begin{array}{ccccccccc}\n2 & 3 & 2 & & 3 & 2 & & 3 & & 7 \\
6 & 7 & 6 & & 7 & 6 & & 7 & 7 \\
8 & 4 & 7 & 6 & & 7 & 6 & & 7 \\
9 & 10 & 7 & 6 & & 7 & 7 & 7 \\
10 & 10 & 10 & & 7 & 7 & 7 & 7 \\
11 & 11 & & 11 & & 7 & 7 & 7 \\
12 & 13 & & 14 & & 7 & 7 & 7 \\
13 & & 14 & & 14 & & 7 & 7 & 7 \\
14 & & 15 & & 7 & & 7 & 7 & 7 \\
16 & & 17 & & 7 & & 7 & 7 & 7 \\
17 & & 18 & & 7 & & 7 & 7 & 7 \\
18 & & 19 & & 7 & & 7 & & 7 & 7 \\
19 & & 19 & & 7 & & 7 & & 7 & 7 \\
10 & & 19 & & 7 & & 7 & & 7 & 7 \\
11 & & 19 & & 7 & & 7 & & 7 & 7 \\
12 & & 19 & & 7 & & 7 & & 7 & 7 \\
13 & & 19 & & 7 & & 7 & & 7 & 7 \\
14 & & 19 & & 7 & & 7 & & 7 & 7 \\
15 & & 7 & & 7 & & 7 & & 7 & 7 \\
16 & & 7 & & 7 & & 7 & & 7 & 7 \\
17 & & 7 & & 7 & & 7 & & 7 & 7\n\end{array}
$$

3.3 C-Property

² the exact C-property, if it is exact when applied to the quiescent flow case.

If the C-property is not satisfied parasitic waves may appear in the numerical solution.

3.4 Test problem A

We wish to consider the propagation of a wave in an enclosed rectangular region $D = f(x; y)$: $x_0 \cdot x \cdot x_M$; $y_0 \cdot y \cdot y_M g$.

For this problem the riverbed is defined as

$$
B(x; y) = \begin{cases} 8 & 3 & 3 \\ \frac{2}{5} & 0.5 & \cos^2 \frac{(x - 0.5) \cancel{u}}{0.25} \quad \cos^2 \frac{(y - 0.5) \cancel{u}}{0.25} \quad \text{if} \quad 0.375 < x < 0.625 \text{;} \\ 0 & \text{and} \quad 0.375 < y < 0.625 \text{;} \\ 0 & \text{otherwise} \end{cases}
$$

which represents a mound positioned in the centre of the region as is illustrated in figure 3.2.

We have the initial conditions

$$
B = \begin{cases}\n 8 & 3 & 3 \\
 \frac{3}{5} & 1 + 0.25 & \cos^2 \frac{(x - 0.225)\frac{y}{A}}{0.25} & \cos^2 \frac{(y - 0.5)\frac{y}{A}}{0.25}\n \end{cases}
$$
\nif $0.1 < x < 0.35$;\nand $0.375 < y < 0.625$ \n
$$
1 \quad \text{B}(x; y) \quad \text{otherwise}
$$
\n
$$
u(x; y; 0) = 0; \quad v(x; y; 0) = 0
$$

Boundary conditions are required at the upstream and downstream boundaries and along the sides of the channel. Since walls are present at all boundaries

Figure 3.2: The riverbed

we need to reflect the velocity

$$
u_{-1j}^n = i u_{0j}^n
$$
; $u_{M+1j}^n = i u_{Mj}^n$; $v_{i,-1}^n = i v_{i,0}^n$; and v^n

Jacobian matrices for the 2D shallow water equations,

$$
\tilde{A} = \frac{\partial F}{\partial w} = \oint_{Q}^{6} \vec{c}^{2} \vec{i} \vec{i} \vec{j} \vec{k} \quad \text{and} \quad \tilde{B} = \frac{\partial G}{\partial w} = \oint_{Q}^{6} \vec{i} \vec{i} \vec{v} \quad \vec{v} \quad \vec{i} \quad \vec{j}
$$
\nwhere $\vec{c} = \frac{1}{g\hbar}$ is the wave speed.
\n
\nThe corresponding eigenvalues and eigenvectors for \tilde{A} are
\n
$$
\begin{vmatrix}\nF = \vec{u} \vec{i} \quad \vec{c} \quad \vec{j} \quad \vec{k} \\
\vec{k} \quad \vec{k} \quad \vec{k} \quad \vec{k}\n\end{vmatrix} = \vec{u} \cdot \vec{i} \quad \vec{k} \quad \vec
$$

an

$$
\tilde{C}_{31} = \tilde{V} \times \tilde{C}_{32} = \tilde{V}_{33} \times \tilde{C}_{33} = \tilde{V} + \tilde{C}_{33}
$$

and

$$
\tilde{\mathbf{e}}_{1}^{G} = \begin{Bmatrix} 2 & 3 & 2 & 3 & 2 \\ 6 & 1 & 7 & 2 \\ 4 & 1 & 7 & 2 \\ 7 & 6 & 3 & 1 \end{Bmatrix}; \quad \tilde{\mathbf{e}}_{2}^{G} = \begin{Bmatrix} 2 & 3 & 2 & 3 \\ 0 & 7 & 7 & 7 \\ 4 & 1 & 7 & 7 \\ 0 & 0 & 7 & 6 \end{Bmatrix}; \quad \tilde{\mathbf{e}}_{3}^{G} = \begin{Bmatrix} 2 & 1 & 7 \\ 6 & 1 & 7 \\ 4 & 1 & 7 \\ 0 & 7 & 7 \end{Bmatrix}.
$$

The Roe averages are given by

$$
\tilde{u} = \frac{P_{\overline{h_R}}u_R + P_{\overline{h_L}}u_L}{P_{\overline{h_R}} + P_{\overline{h_L}}}\;,\quad \tilde{v} = \frac{P_{\overline{h_R}}v_R + P_{\overline{h_L}}v_L}{P_{\overline{h_L}}v_R}
$$

The Coriolis parameter, f, is given by $f = 2 \sin A$ where $= 7.292 \pm 10^{-5} s^{-1}$ is the rotation rate of the earth and \vec{A} is latitude. The source term is then given by

$$
R = f + g = \begin{cases} 2 & 3 & 2 & 3 \\ 6 & 7 & 6 & 7 \\ 6 & 7 & 6 & 7 \\ 7 & 6 & 7 & 7 \\ 8 & 0 & 7 & 7 \\ 0 & 0 & i & 6 \end{cases}
$$

:

For the source term approximations \hat{f} and \hat{g} we use the Roe average values \tilde{u} and \tilde{v} .

3.6.1 The ¯-plane approximation

The use of local cartesian coordinates and the restriction of our attention to only a small region of the globe allows us to model the Earth's surface as a plane, thus avoiding the complications of spherical geometry. However, we can retain the leading order e ect of the Earth's curvature by taking into account the variation of the coriolis parameter, f , with latitude.

The ¯ plane approximation, introduced by C.G Rossby [14], assumes that f varies linearly with the meridional (North-South) distance y. If we let A_0 be the latitude at the origin of our cartesian coordinates, and expand f in a Taylor series about A_0 we get,

$$
f(A_0 + \pm A) = f(A_0) + \pm AF'(A_0) + O(\pm A^2)
$$

for small $\pm \hat{A}$, $y = a \pm \hat{A}$, so that

$$
f = 2 \sin A_0 + \frac{2 \cos A_0}{a} y
$$

$$
= f_0 + y;
$$

where a = 6:37 \pounds 10⁶ is the mean radius of the Earth, and $\bar{a} = \frac{d\pmb{r}}{dy} = \frac{2\Omega\cos A_0}{a}$ $\frac{\cos A_0}{a}$. If the origin of our coordinate system lies on the equator, then $A_0 = 0$, $f_0 = 0$ and

$$
f=\frac{2}{a}y=\sqrt{-y}
$$

3.6.2 Inertial Motion

The coriolis "force"

$$
\mathsf{F}_c = j \ 2 \quad \mathsf{E} \, \mathsf{u}
$$

acts at right angles to the velocity u (and also to) and hence cannot change the speed of a fluid particle. However, it will change the direction of motion.

Consider an open mass of water, such as the ocean. The water is set in motion by an initial impulse, such as a local wind blowing across its surface. It then moves under the influence of the coriolis force and gravity. If there are no other forces acting on the water the movement of the current will be circular.

Now consider a particle moving around a circle of radius R on a plane rotating with angular velocity about the origin. Since the particle travels a distance of $24R$ in time $24=$, its speed of motion U is given by

$$
U = \frac{2\frac{1}{4}R}{2\frac{1}{4}} = R.
$$
 (3.9)

When and R are known, we can use the relationship (3.9) to find the horizontal velocity components u and v . We can specify the curve

$$
U^2 = u^2 + v^2;
$$

by the pair of parametric equations,

$$
u = U \sin ft \tag{3.10}
$$

$$
v = U \cos ft \tag{3.11}
$$

thus giving u and v as functions of time and the Earth's rotation. Such motion is called inertial motion or inertial oscillation.

3.7 Test Problem B

In this problem we wish to simulate the inertial currents observed on the ocean's surface. The domain is a rotating plane, and for simplicity we will assume that the origin of our coordinate system is on the axis of rotation. The seabed is taken to be flat, giving an initially undisturbed surface,

$$
h(x; y; 0) = constant;
$$

with zero initial velocity

 $u(x; y; 0) = 0$; and $v(x; y; 0) = 0$:

The boundaries are prescribed velocities of the form 3.10 and 3.11.

3.8 Numerical Results

3.8.1 Test Problem A

Figures 3.3 show the results of Roe's Q scheme applied to test problem A. The scheme was used with step-sizes $x = y = 0.02$, and $t = 0.002$. For this problem the variable riverbed is given by a cosine squared bell positioned in the centre of the channel. The initial conditions represent an initial pulse on the water's surface which rapidly loses height as it splits and propagates outward in all directions. One of these waves passes over the hump in the riverbed becoming partially reflected. The left moving waves hit the boundary at $x = 0$ and are split further. Since walls are present reflection occurs.

Results are only shown up to a final time of $t = 0.8$, beyond which the waves continue to disperse, rebounding back into the region as they hit the sides of the channel. The surface of the water begins to flatten as the waves gradually become smaller and smaller in amplitude.

It is diecult to evaluate the performance of Roe's Q scheme since no analytical solutions are available for comparison. The results are smooth and appear to be free from any spurious oscillations. However, Roe's scheme is only first order and therefore dissipative. Dissipation occurs when the amplitude of the travelling waves are damped resulting in the numerical solution being smeared. Hence the apparent flattening of the water surface.

3.8.2 Test Problem B

Figures 3.4 show the surface elevation $(h + B)$ for a 100 km $\mathcal E$

case, and their periodic motion is more clearly visible. The initial height of the surface is not maintained and decreases as time increases. At times $t = 90$ & $t = 180$, and $t = 120$ & $t = 240$ (figures 3.7) the wave profiles are almost identical, but the scales to the side of each plot show that the surface height has actually dropped. Again this could be due to the di usive nature of Roe's scheme.

Figure 3.3: Numerical results of Roe's Q Scheme for test problem A - Surface elevation $(h + B)$.

Figure 3.4: Numerical results of Roe's Q Scheme for test problem B - Surface elevation $(h + B)$.

Figure 3.5: Illustration of the e ect of halving the spatial step size from $x = y = 2$ to $x = y = 1$.

Figure 3.7: Numerical results of RoofRohem3(3.72(for(Ro)-btest(of)-4problem(Ro)-bB(of)-4with)]

Chapter 4

Crowd Dynamics

4.1 Introduction

The movement of people within crowds is a complex phenomenon that has been studied for over 40 years. It is of interest in numerous fields such as building development, event management and, particularly in more recent years, emergency planning for evacuation and disaster. Pedestrian movement shares a number of characteristics with tra c flow. Models initially developed to describe tra c flow have successfully been adapted and applied to crowds. Much of the work so far has focused on modelling in one dimension, since tra c naturally moves in single lane formation. Pedestrians, however, tend to inhabit domains where movement is less restricted and can be multidirectional. The aim of this study is to extend the one dimensional model proposed by Payne and Whitham [19] and consider a more realistic two dimensional situation in which individuals can overtake each other and, for example, move across a room towards a doorway.

There are two main approaches to modelling crowds; *microscopic* and macroscopic. Microscopic models consider the individual preferences of each pedestrian and look to capture more of the psychological aspects of human behaviour. They consider how individual pedestrians react to their surroundings, adjusting their behaviour based on the conditions ahead. Dirk Helbing has published many papers in this area (see www.helbing.org for more details). An alternative approach is to consider the overall average behaviour of the crowd as a whole by treating the crowd as a single body with continuous velocity and density fields. This continuum type model is referred to as Macroscopic. Although microscopic models are considered more advantageous in terms of capturing individual sensitivities and interaction behaviour they involve the resolution of a large number of individual parameters and are therefore very expensive. Conversely, macroscopic models are much easier to apply and are generally more computationally e cient.

in the x and y directions respectively, combined with two dynamic equations derived from the Navier Stokes equations of motion for a compressible fluid, but with a pressure $\rho=\,C_0^2\mathscr{Y}$ and with the addition of a relaxation term, thus giving

$$
u_t + u u_x + v u_y = \frac{U(\frac{1}{2}) i}{i} i \frac{C(\frac{1}{2})}{\frac{1}{2}} \frac{1}{2} \tag{4.2}
$$

$$
V_t + UV_x + VV_y = \frac{V(\frac{1}{2}) i V}{\lambda} i \frac{C(\frac{1}{2})}{\frac{1}{2}} \frac{1}{2} y; \qquad (4.3)
$$

where ζ is the relaxation time and $U(\frac{1}{2})$, $V(\frac{1}{2})$ are functions representing the desired velocity of a pedestrian for a given crowd density. $C(\mathcal{B})$ is the anticipation coe cient, which here we take to be $C(\mathcal{V}) = C_0^2$ where C_0 is a constant. For a more detailed discussion on this see [12] & [21].

The relaxation term describes interactions between pedestrians and aims to capture the way in which individuals adjust their velocity towards an optimal value, given by the velocity function. If an individual's speed is greater than this value then the relaxation term has the e ect of slowing the pedestrian down. Conversely, if the individual is moving at a speed less than the desired value the relaxation term speeds the individual up. This change is not instantaneous, the average time required for this process being given by the relaxation time ζ .

This is a much studied model and several variations to it have been suggested. The model has been criticised as under certain conditions it allows negative wave speeds, thus enabling individuals to travel backwards against the fluid flow. In addition, the wave propagation speeds (given by the eigenvalues of the Jacobian matrices of the system 4.4) are such that one will always be greater than the velocity of the fluid flow. This implies that information can travel faster than the flow and reach fluid particles from behind. In terms of tra c modelling, the isotropic nature of the Payne Whitham model contradicts the generally held belief that tra ϵ is an anisotropic fluid. Drivers on a single lane motorway will mainly react to conditions on the road ahead of their vehicle. Although the characteristics of pedestrian interaction are also primarily anisotropic there will be some influence from those behind, particularly in crowds where individuals are moving in close proximity to one another. This model limits us to those interactions that are isotropic assuming that the crowd responds equally to information from behind as well as from in front.

4.2.1 Conservation form

Proceeding in a similar manner to section 3.2 we can rewrite (4.1), (4.2) and (4.3) in the conservative vector form

$$
\begin{array}{ccccccccc}\n2 & 3 & 2 & 3 & 2 & 3 & 2 & 3 \\
& & \frac{1}{2} \\
& & & \frac{1}{2} \\
& & & \frac{1}{2} \\
& & & & \frac{1}{2} & \frac{
$$

4.3 Roe's Q Scheme

The Roe averaged Jacobian matrices of the fluxes are given by

$$
\tilde{A} = \frac{\mathcal{E}}{\mathcal{E}} = \oint_{\tilde{Q}}^{Q} C_0^2 \, i \, \tilde{u}^2 \, 2\tilde{u} \, 0 \, \frac{3}{5} \quad \text{and} \quad B = \frac{\mathcal{E}}{\mathcal{E}} = \oint_{\tilde{Q}}^{Q} \, i \, \tilde{u} \tilde{v} \quad \tilde{v} \, \tilde{u} \, \frac{3}{5} \tilde{u} \tilde{v} \quad \tilde{v} \, \tilde{u} \, \frac{3}{5} \tilde{v} \, \tilde{v} \quad \tilde{v} \, \tilde{u} \, \tilde{v} \,
$$

The corresponding eigenvalues and eigenvectors for \tilde{A} are

$$
\int_{s_1}^{s_1} F = \tilde{u}_1 \ C_0; \quad \int_{s_2}^{s_1} F = \tilde{u}; \quad \int_{s_3}^{s_1} F = \tilde{u} + C_0;
$$

and

$$
\tilde{\mathbf{e}}_1^F = \begin{matrix} 2 & 3 & 2 & 3 & 2 & 3 \\ 2 & 1 & 7 & 6 & 7 \\ 4 & 7 & 7 & 6 & 7 \\ 6 & 1 & 7 & 7 & 7 \end{matrix}; \quad \tilde{\mathbf{e}}_2^F = \begin{matrix} 2 & 3 & 2 & 3 \\ 2 & 7 & 7 & 7 \\ 2 & 0 & 7 & 7 \\ 5 & 6 & 8 & 7 \end{matrix}; \quad \tilde{\mathbf{e}}_3^F = \begin{matrix} 2 & 1 & 7 \\ 2 & 7 & 7 \\ 4 & 7 & 7 \\ 6 & 7 & 7 \end{matrix}
$$

For matrix \tilde{B} we have

$$
\tilde{G} = \tilde{V}j \quad C_0; \quad \tilde{G} = \tilde{V}; \quad \tilde{G} = \tilde{V} = \tilde{V}
$$

and

$$
\tilde{\mathbf{e}}_{1}^{G} = \begin{matrix} 2 & 3 & 2 & 3 & 2 & 3 \\ 6 & 1 & 7 & 6 & 0 & 7 \\ 7 & 7 & 7 & 6 & 7 \\ 8 & 6 & 7 & 7 & 6 \\ 1 & 1 & 1 & 1 & 1 \\ 1 & 1 & 1 & 1 &
$$

The Roe averages are given by

$$
\tilde{u} = \frac{P_{\overline{\mathcal{H}}_{R}} u_{R} + P_{\overline{\mathcal{H}}_{L}} u_{L}}{P_{\overline{\mathcal{H}}_{R}} + P_{\overline{\mathcal{H}}_{L}}}, \quad \tilde{v} = \frac{P_{\overline{\mathcal{H}}_{R}} v_{R} + P_{\overline{\mathcal{H}}_{L}} v_{L}}{P_{\overline{\mathcal{H}}_{R}} + P_{\overline{\mathcal{H}}_{L}}}
$$
 and $\tilde{v}_{2} = P_{\overline{\mathcal{H}}_{R}\overline{\mathcal{H}}_{L}}.$

4.4 The Velocity Function

The exact form of the velocity function is dependent on the nature of the physical situation and as a consequence many choices have been proposed, see [12]. A pedestrians movement is guided by his or her desired velocity which is composed of a desired speed and a desired direction. It is assumed that desired speed has mean 1:34 ms^{-1} and standard deviation 0:26 ms^{-1}

[10]. The desired direction of movement is given by the pedestrian's intended destination.

4.4.1 The Bee-line E ect

In this report we consider the bee-line e ect where pedestrians wish to move towards a specific target, such as an exit, positioned at certain point in the room. The pedestrians aim is to reach the point by taking the most

4.5 Initial and Boundary Conditions

To simulate the bee-line e ect we need to first specify the initial and boundary conditions. We wish to consider a crowd in a rectangular room with solid walls. There is an exit along the wall $x = x_M$ towards which the crowd wishes to head. The initial conditions are given by a gaussian distribution of the form

$$
V_2(x; y; 0) = V_0 \exp \frac{1}{i} \frac{(x i \partial)^2}{c^2} i \frac{(y i \partial)^2}{d^2}.
$$

and the crowd is initially at rest, i.e.

$$
u(x; y; 0) = 0;
$$
 and $v(x; y; 0) = 0.$

The question remains as to what to do at the boundaries. The solid walls of the room can be reproduced by imposing the slip boundary condition

$$
u \mathrel{\ell} n = 0;
$$

i.e. the pedestrians cannot walk through the walls but can move tangentially to*j*thesitJ/F411.9T-025(is7.89211.i;N11.95.324.74T)2.81 EQuTDE7d1TD[(721.79TD[i5e6976e7/F211.95Tf6.66ITD[(3.1T2[(x6ITD[(3.1T2[(x6IT.95Tfn26(ther .guTDE7d1TD[(721.79TD[i5e6976e7/F211.95Tf6.66ITD[(3.1T2[(x6ITD[(3.1T2[(x6IT.95Tfn26(ther However, we need to ensure that pedestrians do not re-enter the room via the exit. If the incorrect boundary conditions are used then the pedestrians may be reflected back into the domain resulting in an inaccurate numerical solution. Transmissive boundary conditions can be used to eliminate the left moving waves (pedestrians) but allow the right moving waves to pass through the boundary, thus enabling the crowd to leave the room without being reflected back. One type of transmissive boundary condition discussed by Godlewski and Raviart [5] can be obtained by using the k -Riemann invariants, given by

$$
\frac{\varpi r_k}{\varpi \mathbf{u}} \mathbf{\ell} \mathbf{e}_k^F = 0 \tag{4.5}
$$

and

$$
\frac{\mathscr{Q}S_k}{\mathscr{Q}\mathbf{u}}\,\mathscr{E}\mathbf{e}_k^G=0\tag{4.6}
$$

for the x and y sweeps respectively. ${\sf e}^F_k$ and ${\sf e}^G_k$ are the eigenvectors of the Jacobian matrices **A** and **B**, and the r_k , s_k denote the *k*-Riemann invariants where k is the k^{th} component of the system. For the system (4.4) we have $k = 3$.

Since the exit lies along the boundary at $x = x_M$, we are only concerned with the direction of the u component of the pedestrians velocity and therefore need only consider solutions to (4.5), i.e.

$$
\frac{\varpi_{1}}{\varpi_{2}} + (u_{j} C_{0}) \frac{\varpi_{1}}{\varpi_{2}} + v \frac{\varpi_{1}}{\varpi_{2}} = 0
$$
\n
$$
C_{0} \frac{\varpi_{2}}{\varpi_{2}} = 0
$$
\n
$$
\frac{\varpi_{3}}{\varpi_{2}} + (u + C_{0}) \frac{\varpi_{3}}{\varpi_{2}} + v \frac{\varpi_{3}}{\varpi_{2}} = 0
$$

These equations can be solved to yield

$$
r_1 = u_j
$$
 $C_0 \ln \frac{1}{2}$; $r_2 = u$; $r_3 = u + C_0 \ln \frac{1}{2}$.

Hence we obtain the following approximations of u and $\n *k*$,

$$
u = \frac{1}{3}(r_1 + r_2 + r_3)
$$
; and $k = \exp \frac{1 - r_3}{2C_0}$.

We can use these approximations at the exit to ensure that outgoing waves pass through and are not artificially reflected back. In the case $\sqrt{1} < 0$, $\frac{1}{2}$ $\frac{1}{2}$ > 0, $\frac{1}{3}$ > 0, r_1 represents the left moving waves and r_2 and r_3 the right moving waves. Thus we want to eliminate r_1 but allow r_2 and r_3 to pass through the boundary. This can be done by setting

 $r_1 = u^*_{M;j} i$ C₀ In $\mathbb{Z}_{M;j}^*$; **U**+

4.6 Numerical Results

A series of numerical experiments were used to calibrate the model with the aim of determining an optimal set of parameter values that produced behaviour as close as possible to what we would physically expect. This involved varying the size of the room, the starting position of the crowd and the location of the exit. It was essentially a trial and error method but produced some intriguing results. In particular, it was found that the model was extremely sensitive to changes in the value of the anticipation coe cient C_0 .

Figures 4.1 and 4.2 show the results for a $10mE10m$ room, with $C_0 = 0.4$, spatial step sizes $x = y = 0.2$ and a time step $t = 0.02$. Initially the crowd is gathered around the point $(3/3)$ and the exit is positioned in the top right hand corner of the room between the points $y = 7$ and $y = 9$. The crowd remains tightly packed and slowly moves across the room reaching the exit at around $t = 10$. The density builds creating a pile up across the doorway, the crowd is unable to leave the room and density increases even further. The code runs up until a time of $t = 14.5$, after which problems arise at the boundary where unrealistically high density values mean that the scheme becomes unstable.

Figures 4.3, 4.4 and 4.5 show the e ect of doubling the size of the anticipation coe cient to $C_0 = 0.8$. The same initial data, spatial and temporal step sizes were used. The behaviour of the crowd becomes very strange. Unlike in the above example, the crowd is much less densely packed and therefore moves more freely. After 4 seconds the shape of the crowd starts to become distorted and appears to move away from the exit towards the

the room.

In general, the smaller the value of C_0 the greater the clustering e ect and the higher the density. It appears that when $C_0 = 0.4$ the crowd moves too slowly through the doorway. Therefore a blockage occurs further restricting their movement and preventing an e cient exit. In section 4.2 we stated that the anticipation coe cient is related to pressure. If C_0 is low then the pressure is low and the magnitude of the wave speeds $\binom{1}{3}$, $\binom{3}{3}$ are smaller. The slow speed of movement means that the crowd remains closely assembled and thus too many pedestrians reach the exit at the same time. Conversely, when C_0 is high the pressure is high and the waves travel much faster, dispersing the crowd and enabling the pedestrians to exit the room in an e cient manner. To maintain stability and prevent spurious solutions, these speeds need to be controlled via the time step t . One possibility would be to employ a variable time step. It appears from trials so far that when C_0 = constant there does not exist a value for which the two e ects can be balanced. This suggests that a density dependent anticipation coe cient would perhaps be more adequate. The observed distortions may also be in part due to the dimensional splitting. A more e ective scheme would be fully two dimensional.

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Figure 4.1: 2D crowd model - $10mE10m$ room with exit at $x = 10, 7 \cdot y \cdot 9$, $C_0 = 0.4$, $x = y = 0.2$, $t = 0.02$.

Figure 4.2: 2D crowd model - $10mE10m$ room with exit at $x = 10, 7 \cdot y \cdot 9$, $C_0 = 0.4$, $x = y = 0.2$, $t = 0.02$.

Figure 4.3: 2D crowd model - $10mE10m$ room with exit at $x = 10, 7 \cdot y \cdot 9$, $C_0 = 0.8$, $x = y = 0.2$, $t = 0.02$.

Figure 4.4: 2D crowd model - $10mE10m$ room with exit at $x = 10, 7 \cdot y \cdot 9$, $C_0 = 0.8$, $x = y = 0.2$, $t = 0.02$.

Figure 4.5: 2D crowd model - $10mE10m$ room with exit at $x = 10, 7 \cdot y \cdot 9$, $C_0 = 0.8$, $x = y = 0.2$, $t = 0.02$.

Figure 4.6: 2D crowd model - $10mE10m$ room with exit $x = 10$, $2 \cdot y \cdot 4$, $C_0 = 0.4$, $x = y = 0.2$, $t = 0.02$.

Figure 4.7: 2D crowd model - $10mE10m$ room with exit at $x = 10$, $2 \cdot y \cdot 4$, $C_0 = 0.8$, $x = y = 0.2$, $t = 0.02$.

Figure 4.9: 2D crowd model - $10mE10m$ room with exit at $x = 10, 2 \cdot y \cdot 4$, $C_0 = 1.3$, $x = y = 0.2$, $t = 0.02$.

Chapter 5

Conclusions and Further Work

The aim of this dissertation was to develop a numerical method for approximating two dimensional systems of hyperbolic conservation laws with source terms and then apply it to the variable depth shallow water equations and a macroscopic model of crowd flow. The technique of dimensional splitting was used to decompose the system of conservation laws into two parts that could be solved alternately in the x and y directions. We then adopted the Q scheme method of Bermúdez & Vázquez, using a two dimensional extension of Roe's Q scheme to numerically approximate the equations.

In chapter 3 we looked at the application of our method to the variable depth shallow water equations, first using a non-rotating Eulerian frame of reference and then extending the equations to include the e ect of the Earth's rotation. This was done by adding a coriolis term to the source. The numerical solutions produced by the scheme were smooth and oscillation free. However, it was noted that the absence of any exact solutions made it diefcult to fully appraise the accuracy of the method. We observed that the scheme su ered from dissipation leading to smearing of our numerical solution over

time. This could be lessened somewhat by refining the computational mesh but this consequently reduced the e ciency of the code. It was suggested that a better way of eliminating this di usion would be to introduce a flux limiter which switches between a second order approximation when the region is smooth and a first order approximation near a discontinuity.

In chapter 4 we developed a two dimensional model of crowd flow from a one dimensional tra $\,$ c flow model proposed by Payne and Whitham. A series of numerical trials were then used to evaluate the performance of the model. We then sought to interpret the resulting behaviour in terms of the pressure and wave speeds. It was found that the model was extremely sensitive to the value of the anticipation coe cient C_0 . Under certain conditions the model did exhibit some attributes of crowd flow but a number of curious e ects were produced. Computational errors were also experienced caused by the generation of unphysical density values, particularly near the boundaries.

Crowd dynamics is a vast and complex field and in this report we have barely touched on one aspect. The problem is not only multi-dimensional but also multi-faceted. The Payne Whitham model was originally developed for tra c and more work needs to be done to tune it to pedestrian flow. However, the approach is very simplistic and will never capture all the complex interactions required for a more realistic crowd flow model. In terms of the work undertaken so far, there is much room for improvement. Further analysis is required on stability, and finding a suitable anticipation coe cient C_0 that prevents the generation of unphysical density values and yields properties more akin to crowd flow. More generally, there remain numerous possibilities for further investigation, from simple changes to the

initial data and the structure of the domain of interest, such as including an entrance or multiple exits, to more advanced refinements, such as using a microscopic approach to allow for the psychological behaviour of humans, modelling individual preferences and physical interactions.

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