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The Assimilation of Satellite Derived Sea Surface Temperatures into a Diurnal Cycle Model

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NUMERICAL ANALYSIS REPORT 8/2007

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sceptible to discern arising signs, therefore the timing of the observations
and the conditions in which

tion comparisons reflect errors only for different regions and time scales. In diurnal cycle modeling the high winds are not of concern so diurnal sign for the high wind speeds however the diurnal timing is very sensitive to slight changes in wind speeds that the ocean sees. The wind speed w is important because wind stress increases roughly as w and wind speed increases with w therefore the insight arises in NWP sea surface winds speeds can lead to systematic errors in ocean circulation models that are forced by these winds. Wind speeds of less than 5 m s^{-1} account for nearly half of global hourly averaged winds. These winds are concentrated in the tropics and

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here I is the tot s rf ce so r r di tion nd $I_{\downarrow}$ is the s rf ce inso tion nd cer
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$$n = \frac{\beta \int_T^{T_{\downarrow}} I_{\downarrow} - \alpha \int_T^{T_{\downarrow}} I \, dt}{\int_T^{T_{\downarrow}} I_{\downarrow} - \alpha \int_T^{T_{\downarrow}} I \, dt}.$$

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This section describes a method for tuning the code co-er and the surface wind speed parameters as they appear in the forcing algorithms for the GOM mode. It is assumed that these two parameters rely to the most poorly on the N/P or reanalysis mode output. The parameters are used to help GOM fit the site's residuals including the diurnal cycle component.

3.1 The Algorithm

If we first consider the mode ed θ as a function of the fraction code co-er n and wind speed forcing $w = \sqrt{u_{ref} v}$

$$\theta = \theta(n, w). \quad (3.1)$$

Parameters ϵ_A and ϵ_B are introduced here

$$n = n_{ref} + \epsilon_A, \quad (3.2)$$

$$w = \epsilon_B w_{ref}, \quad (3.3)$$

$$\epsilon_B = \sqrt{u_{ref} v}. \quad (3.4)$$

These dependent parameters are assumed to remain fixed over each hour time window. Although the background $u_{ref} v$ and n_{ref} from ECM Forecast every six hours, the code correction is seen to be a small error. Here the wind correction is a fraction error. This avoids corrections to wind direction and allows the strict limits on code co-er $-1 \leq n \leq 1$ to be satisfied. The function can now be written as a function of the parameters

$$\theta = \theta(\epsilon_A, \epsilon_B). \quad (3.5)$$

We now define a cost function $J = J(\epsilon_A, \epsilon_B)$ as

$$J = \sum_{i=1}^N (H(\theta^{model}_i - \theta^{obs}_i))^2, \quad (3.6)$$

where N is the number of observations over the hour window and $t = Td + t$ $Td = T$

$\epsilon_B > \epsilon_A$ the distributed system problem is not solved for ϵ_A

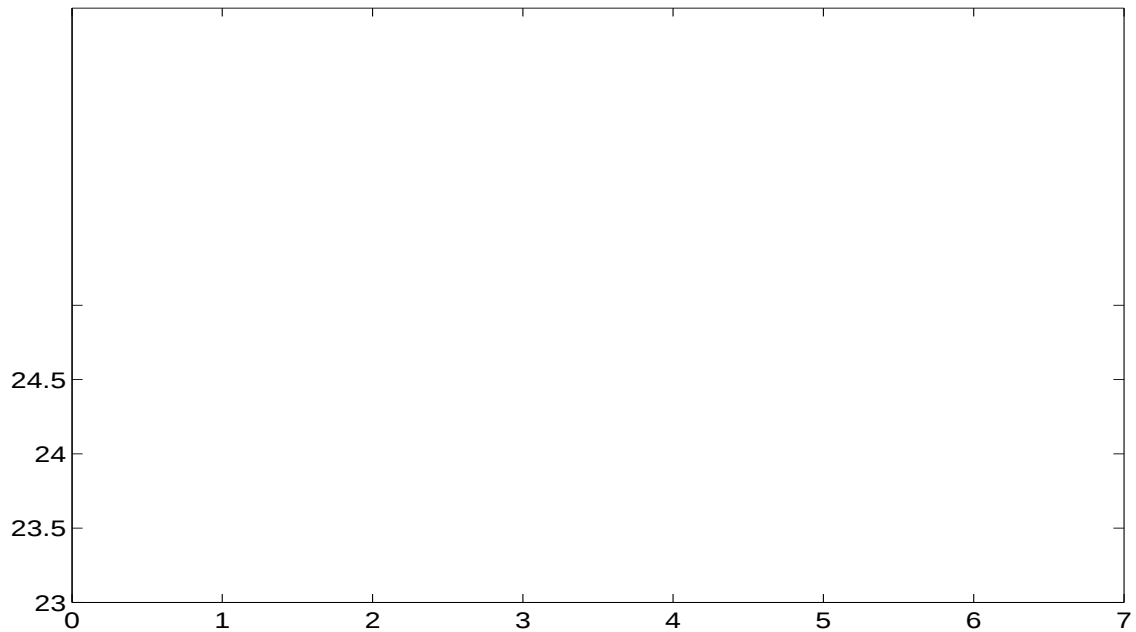
An 'optimal' parameter pair $(\epsilon_A^*, \epsilon_B^*)$ is sought such that for all feasible (ϵ_A, ϵ_B)

$$|J(\epsilon_A^*, \epsilon_B^*)| \leq |J(\epsilon_A, \epsilon_B)|.$$

It is possible that an increase in code coherence and decrease in wind speeds and ice drifts could provide the desired effect in this scenario

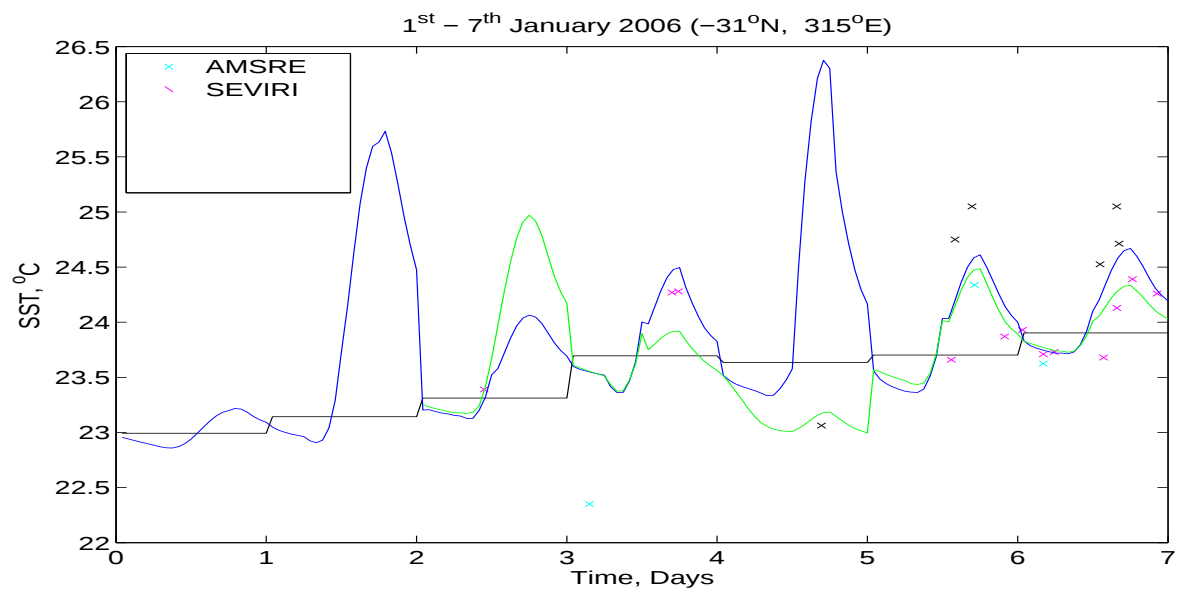
o ser tions fro AM R E nd M if oth types R nd M of o ser tions
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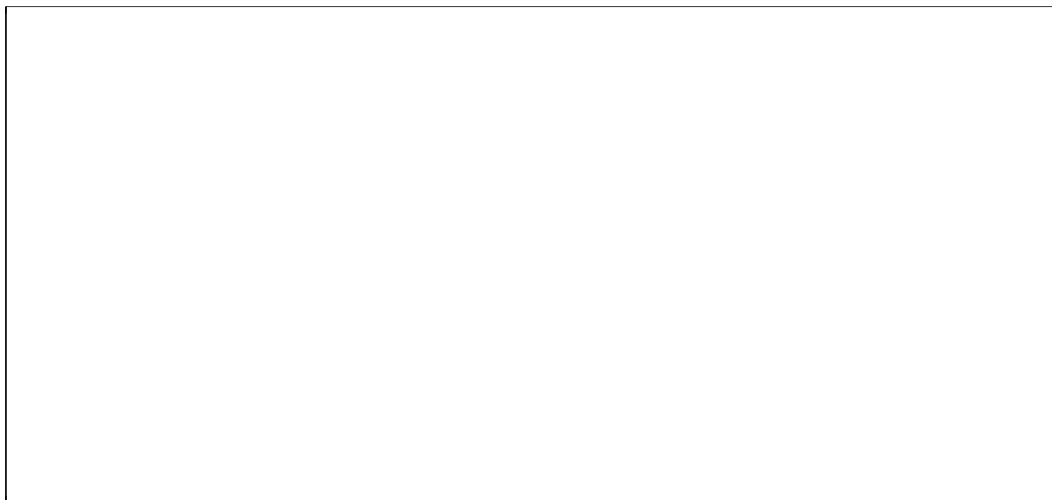
to underestimates the forcing on these occasions. However, earlier in the year the modelled diurnal forcing estimates are larger than the observations suggest. The diurnal assimilation method reduces the model if necessary in the presence of E(R) observations for a daily correction to the wind speed forcing and then the model fractions when E(R) observations are not present. The assimilation is able to reduce the modelled forcing for days through to and increase the diurnal forcing on days and thus fitting the observations much more closely. On days and the assimilation has not been able to reduce the forcing such as the observations would suggest. In these cases the system does not fully adhere to the assumptions of the assimilation routine either because the estimated sensitivity over the parameter range is incorrect or the required change in forcing is outside the stated restrictions.

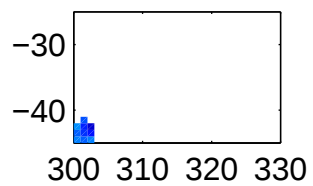


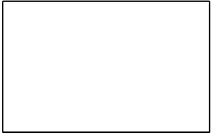
day. Of (A shows that this is compared to the observations and this is why the case of the simulation runs to find this point.

The simulation is essential if correction to the diurnal cycle is based on single observation particularly if it occurs early in the morning. This can be seen in Figure 4 where on day 5 slight correction early in the morning phase leads to larger diurnal cycle than that not tested by further observations and on day 6 there is a result of possible erroneous cycloidal observation strongly related to evening in the diurnal cycle. These examples illustrate how the scheme could be further improved in the future by incorporating more observations and finding on the noise edge gained to forecast the trend of observation error both systematic and random within the simulation cycle.









4.2 Comparing Different Satellite Observations

Further analyses were so performed to assess the errors associated with individual observation types relative to the control and these are presented in Table 4.2.

	No. Obs	Mean	STD	RMS
GOM E (R)	12	-0.1	0.1	0.1
GOM AM RE	12	0.1	0.1	0.1
GOM M	12	0.1	0.1	0.1

Table 4.2 Results showing the number of observations, the mean, STD, and RMS of $\theta_{\text{control}} - \theta_{\text{obs}}$, in $^{\circ}\text{C}$, for individual satellite types for the area 10°N to 10°N and 10°E to 10°E during 1st to 10th January 2006.

A similar number of E (R) and AM RE observations were collected over the time period in this region with slightly fewer M observations. The mode of observation techniques reveal differences between the three satellite instruments. The E (R) observations are shown to be on average further than the parameterised surface temperature here as the AM RE and M observations are closer on average than the mode of observation. This suggests that the observations have some systematic errors in this region at this time with E (R) systematic errors too large and/or AM RE and M observations systematic errors too close. The mode could so have been estimated too great a correction. This seems to imply the parameterised corrections in correction for this period are on average 0.1°C less than the E (R) on average difference. The mode of observations are dependent on the O (A) at the start of the day therefore any errors in O (A) will so appear as a correction. The STD and RMS are significantly lower when comparing E (R) observations with either AM RE or M. The largest errors are found with the M observations where the RMS error approaches 0.1°C .

4.3 Day-Night Comparisons

Differences in night time between the restricted hours of observation and daytime between the hours of observation techniques were so observed so

for heighten di rn cyc e nd dr s the ode c oser to the o ser tions the D
 nd RM di erences re in si i r for oth d y nd night

	Me n	D	RM
Before Assi i tion			
d yti e			
night ti e			
After Assi i tion			
d yti e			
night ti e			

Figure 4.4. Results showing the mean, STD, and RMS of $\theta_{\text{model}} - \theta_{\text{obs}}$, before and after the assimilation, in °C, during daytime (10{16) and night time (22{04) local time for the area 10°N to 20°N and 100°E to 120°E during 1st{ 5th January 2006.

4.4 Comparisons to OSTIA

To help determine to what extent the biases

re found to be larger than night time values and when comparing observations
 of -1.5°C is found. This indicates that the site observations on average
 are warmer than OJA's so do not expect to see OJA is largely restricted
 to night time. The sharp difference in day and night time values demonstrates
 the presence of diurnal signals in the daily observations. OJA is the mean of
 of these observations since it is others and so the expectation is that the
 values in this time period observations are increased. Here OJA is forced
 by eliminating daily observations then with wind speeds less than 5 m s^{-1} . These
 additional observations are therefore contributing to the slight cooling in

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